



# Snow Cover Dynamics and Snowmelt Runoff Products & Services - Indian Himalayan River Basins under NHP

Project Director:Dr. V.V.Rao, DD(RSAA)Deputy Project Director:P.V.Raju, Group Director (WRG)

PI: B.Simhadri Rao, Sc/Er"SG", WRAD Co-PI: C.Sai Krishna, Sc/Er"SD", WRAD

Research Scientist: Gopi Senior Research Fellow: V.Madhavi

Water Resources Group National Remote Sensing Centre Indian Space Research Organisation



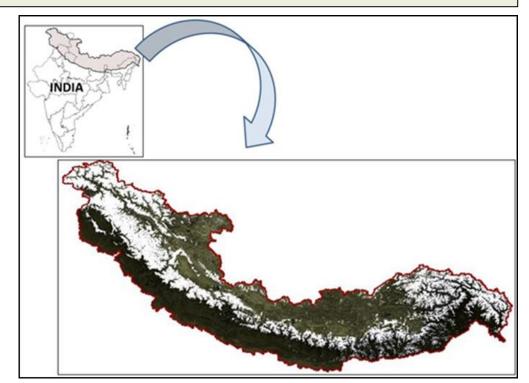


#### **Study Objectives**

- To develop a short-term snowmelt runoff forecasting model using satellite derived products and field data
- To generate a spatial daily gridded snowmelt product
- To generate a spatial 3-day snowmelt forecast gridded product
- To provide short term snowmelt runoff forecast at selected basin outlets during snowmelt season

#### **Study Area**

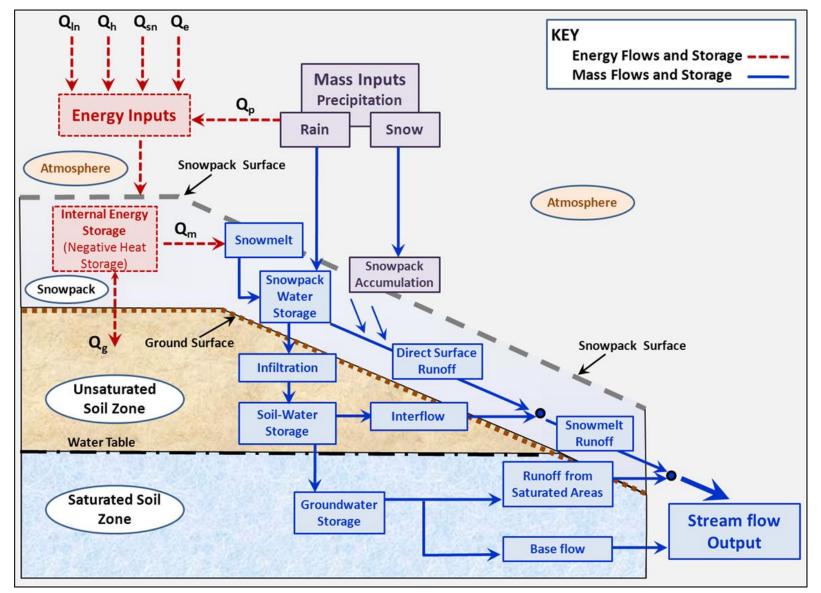
 Indian Himalayas covering Major river systems (Indus, Ganga and Brahmaputra) including outside Indian boundary covering 9.89 Lakh sq.km.





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#### **Schematic of the Snowmelt Process**



Source: Price, Hendrie, and Dunne (1979)



#### **Snowmelt Runoff Estimation**

□ Temperature Index method

Energy Budget method

**Energy Budget for Snowmelt** 

The energy available for snowmelt

$$Q_{m} = Q_{sn} + Q_{ln} + Q_{h} + Q_{e} + Q_{g} + Q_{p} + Q_{i}$$

- Q<sub>sn</sub> = Shortwave net radiation
- Q<sub>In</sub> = Longwave net radiation
- $Q_h$  = Sensible heat flux
- $Q_e$  = Latent heat flux
- $Q_g$  = Ground heat conduction
- $Q_p$  = Energy contained in the rainfall

Q<sub>i</sub> = Change in internal energy

The amount of snowmelt

$$M = \frac{Q_m}{L_f \rho_w B}$$

$$\begin{split} M &= \text{Snowmelt, m} \\ Q_m &= \text{flux density of melt energy, J/m^{-2}day} \\ L_f &= \text{latent heat of fusion, Jkg^{-1}} \\ 0.334X10^6 \text{ Jkg^{-1} at } 0^0 \text{ C} \\ \rho_w &= \text{density of liquid water, Kgm^{-3}} \\ B &= \text{Thermal quality of snowpack} \end{split}$$



#### Energy Components for Snowmelt

Shortwave net radiation

 $Q_{sn} = (1-\alpha) R_s$ 

#### Longwave net radiation

$$Q_{ln} = L\psi - L\uparrow$$
$$= L\psi - \{\epsilon_s \sigma T_s^4 + (1-\epsilon)L\psi \}$$
$$L\psi = \epsilon_a \sigma T_a^4$$

Sensible heat flux

$$\mathbf{Q}_{h} = \rho_{a} c_{p} C_{h} u_{a} (\mathbf{T}_{a} - \mathbf{T}_{s})$$

Latent heat flux

 $Q_e = [\rho_a 0.622L/P_a]C_e u_a (e_a - e_s)$ 

- $\alpha$  snow albedo
- ${\rm R}_{\rm s}$   $\,$  Incoming solar radiation  $\,$
- $L\psi$  Incoming longwave radiation
- T<sub>s</sub> Snow surface temperature
- $\boldsymbol{\epsilon}_s~$  Emissivity of snow surface
- $\boldsymbol{\epsilon}_{a}~$  Emissivity of atmosphere
- T<sub>a</sub> Air temperature
- $\rho_a$  Density of air
- c<sub>p</sub> Specific heat of air
- $C_{h}$  Bulk transfer coefficient for sensible heat
- $\ddot{u_a}$  Wind speed
- L Latent heat of vaporisation
- P<sub>a</sub> Atmospheric pressure
- $\mathbf{C}_{\mathrm{e}}~$  Bulk transfer coefficient for latent heat
- $\tilde{e_a}$  Atmospheric vapour pressure
- $\mathbf{e}_{s}$  Vapour pressure at the snowpack surface



#### **Energy Components for Snowmelt**

Ground heat conduction

 $Q_{g} = k_{g}(T_{g}-T_{sb})/(z_{2}-z_{1})$ 

Energy contained in rainfall

 $Q_{p} = P_{r}\rho_{w}c_{w}(T_{r}-T_{s})$ 

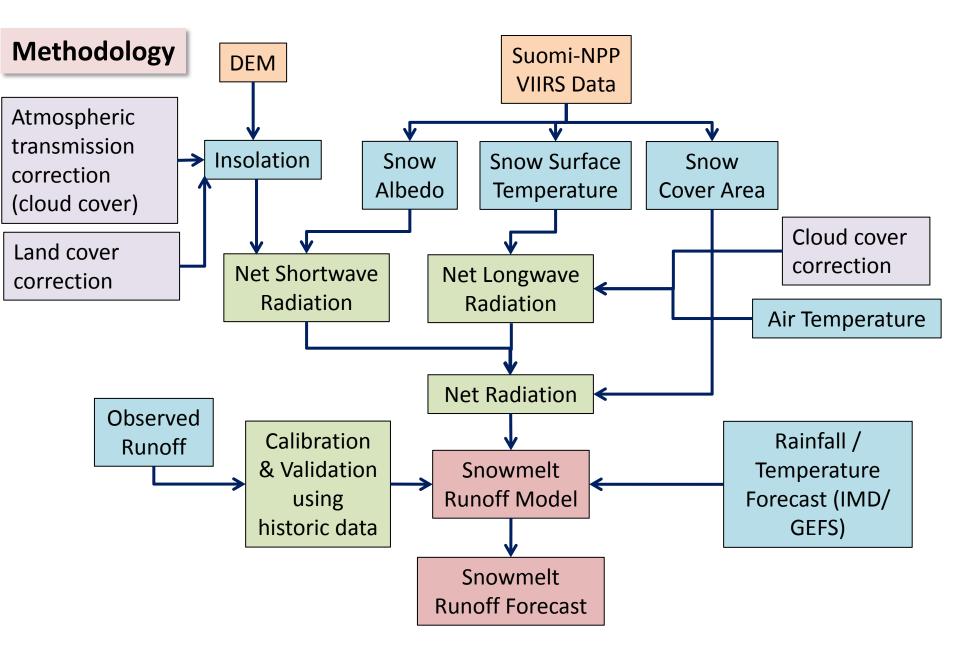
Change in internal energy

 $Q_{h} = [\rho_{s}c_{i}](\Delta T_{s}/\Delta t)\Delta z$ 

- Thermal conductivity of soil Kσ
- Soil depth 7
- Soil temperature at depth  $z_2$
- T<sub>g</sub> T<sub>sb</sub> Temperature at base of snowpack depth  $z_2$
- **Rainfall intensity** P,
- Density of liquid water  $\rho_w$
- Specific heat of liquid water C.,,/
- Temperature of rain (assumed as  $T_a$ ) T,
- Density of snowpack ρ
- Specific heat of ice Ci
- time t
- Height above ground (z=0 at ground; 7 z=d at snowpack surface)
- Snowpack depth d



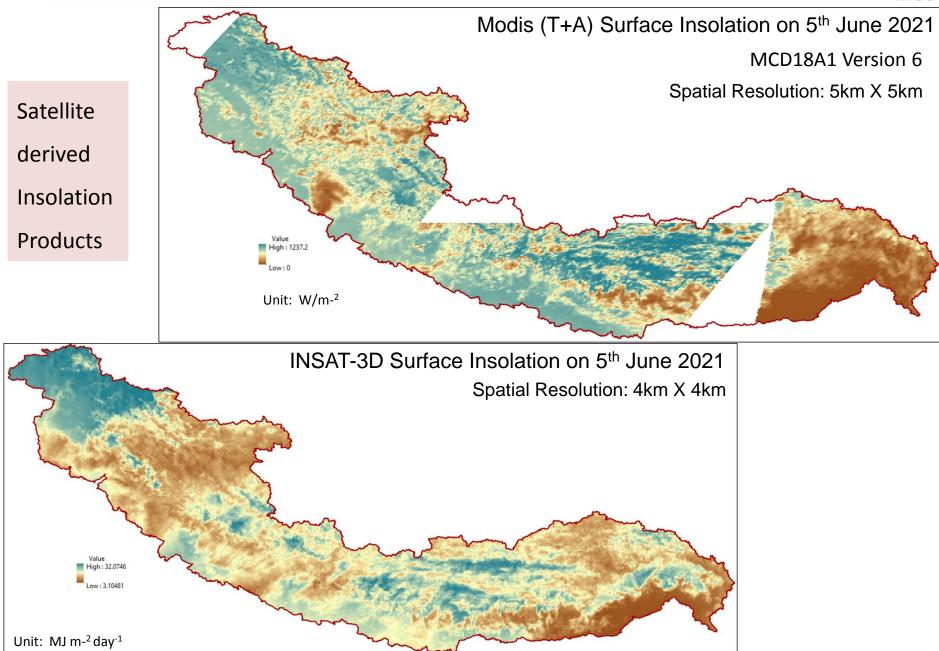




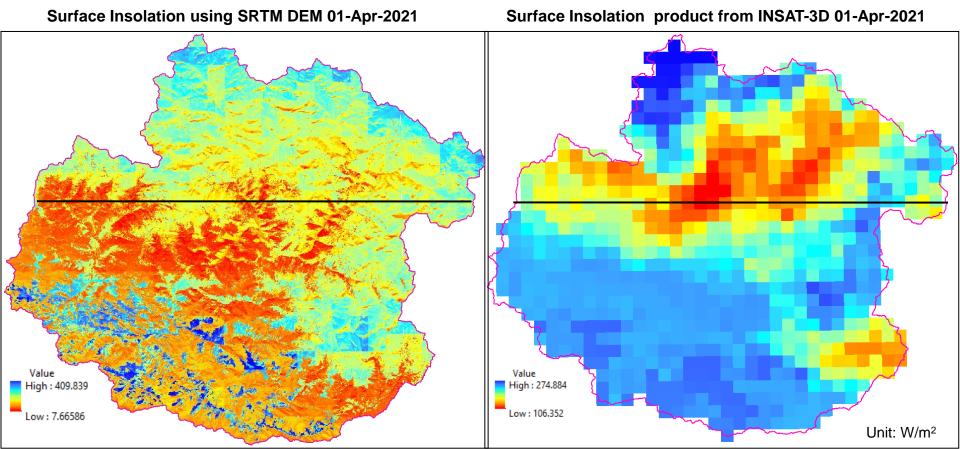
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## **Snow Cover Dynamics and Snowmelt Runoff Products & Services**

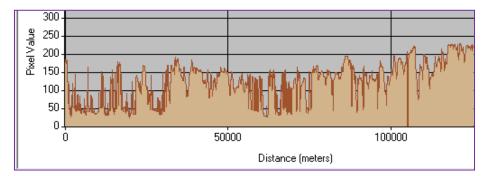
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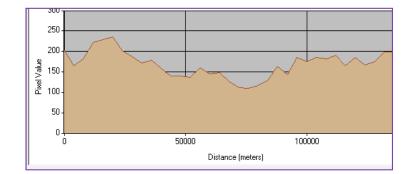




Spatial Resolution:90m



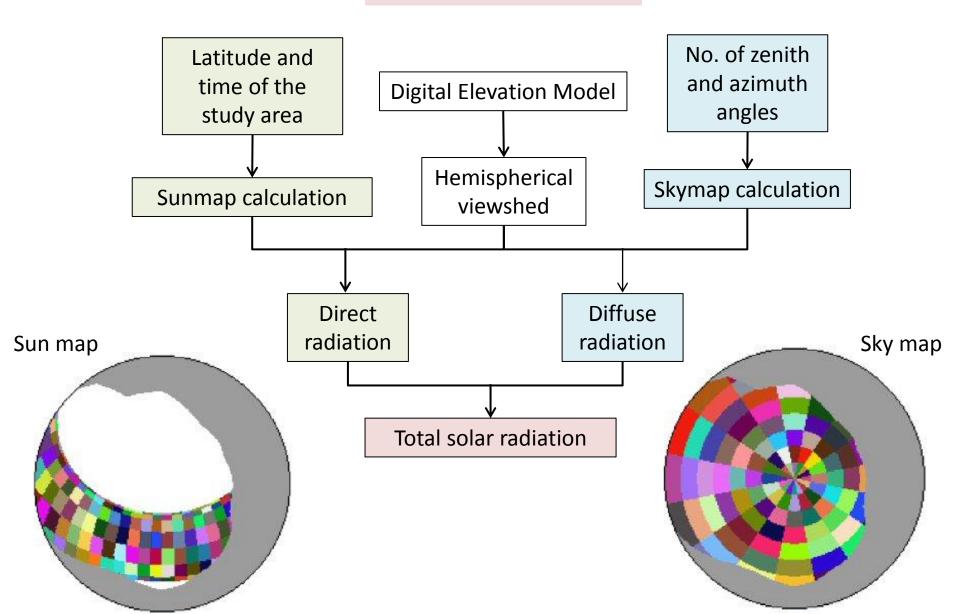
Spatial Resolution: 4 km





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#### Computation of Insolation







#### **Computation of Insolation**

#### <u>Methodology</u>

• Direct radiation is estimated by overlaying viewshed on direct sunmap

 $R_{dir} = S_{const}^* \beta^{m(\theta)} SunDur^*SunGap^*cos(AngIn)$ 

Where

 $S_{const}$  = Solar constant;  $\beta$  = transmissivity of the atmosphere; m( $\theta$ ) = relative optical path length;  $\theta$ =solar zenith angle SunDur = time duration represented by the sky sector; SunGap = gap fraction for the sunmap factor Angln = angle of incidence between the centroid of the sky sector and axis normal to the surface

• Diffuse radiation is estimated by overlaying viewshed on diffuse map

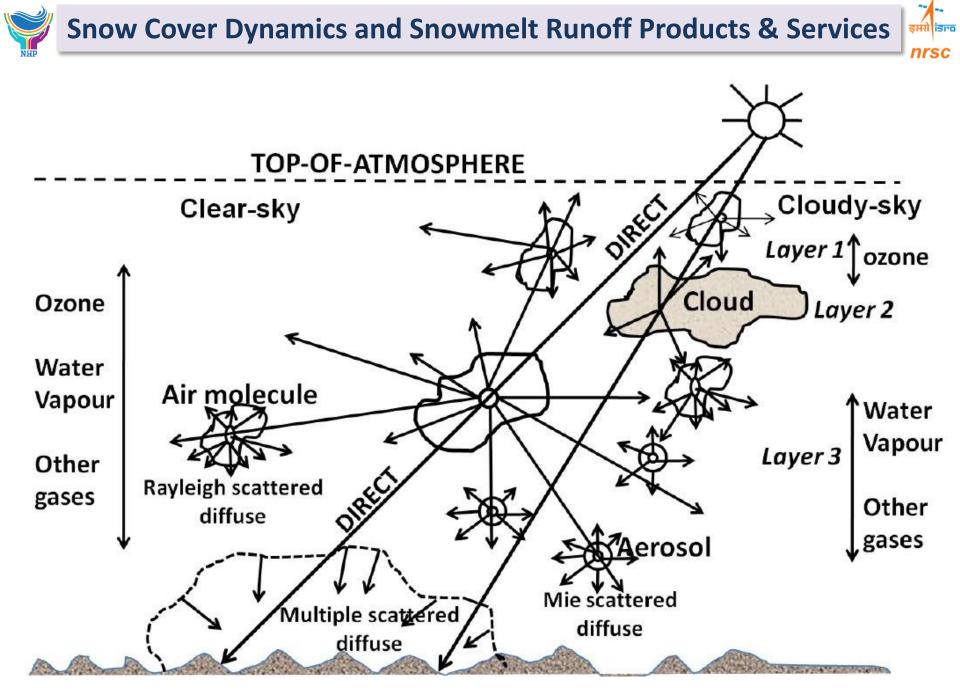
R<sub>dif</sub> = R<sub>glb</sub>\*P<sub>dif</sub>\*Dur\*SunGap\*Weight\*cos(AngIn)

Where

R<sub>glb</sub> = global normal radiation; P<sub>dif</sub> = proportion of global normal radiation that is diffused; Dur = time interval for analysis Weight= proportion of diffuse radiation originating in a given sky sector and the intercepting surface

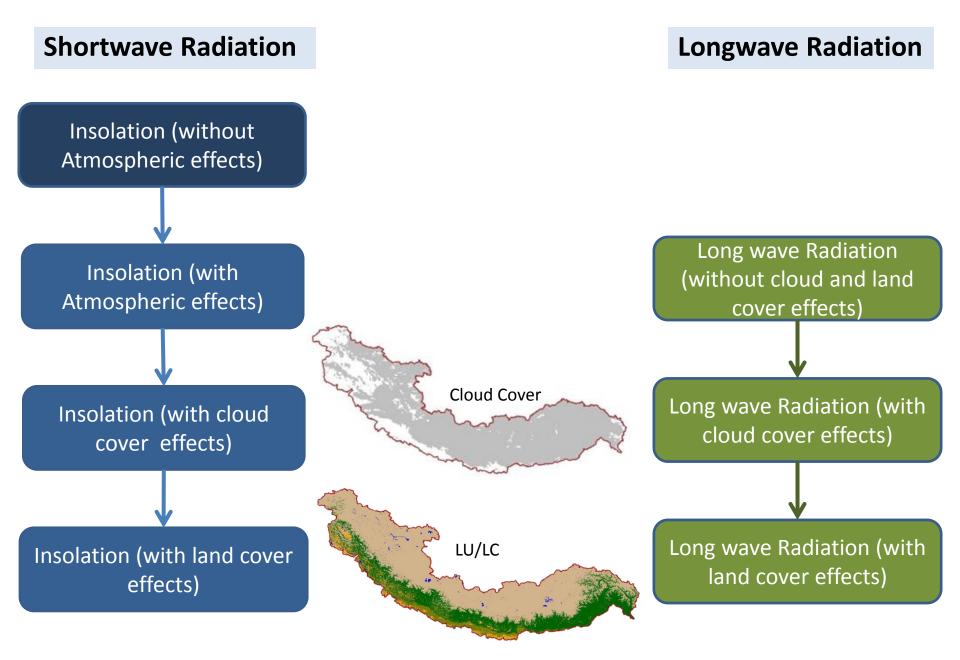
• The total solar radiation is estimated by addition of direct and diffuse radiation

$$R_{global} = R_{dir} + R_{dir}$$



Source: INSAT ATBD document(2015)

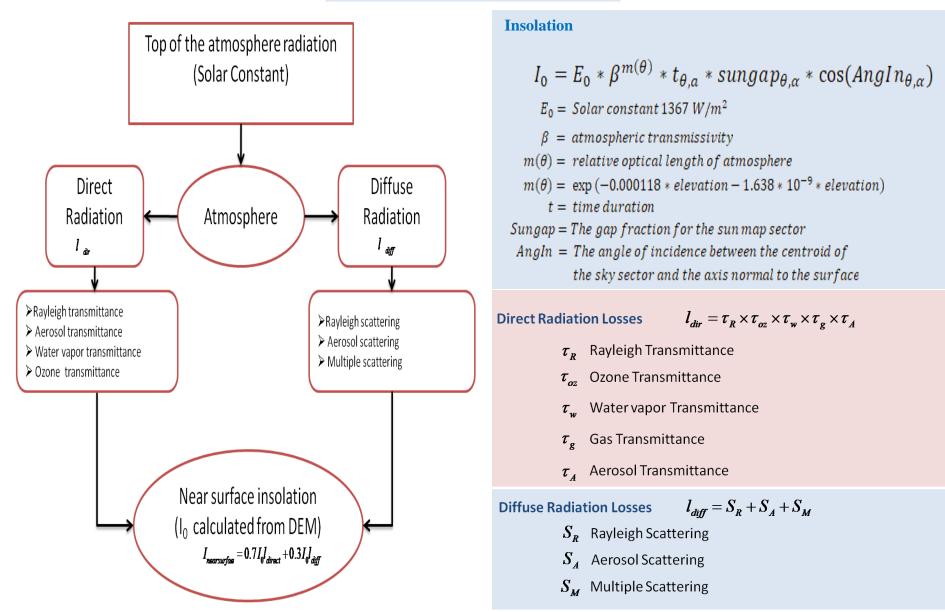








#### **Shortwave Radiation**





#### **Transmission Losses Through Direct and Diffuse Radiation**

>Water vapor, Ozone, Aerosol Transmittance

$$\tau(x) = \exp[-x(a+bx+cx^d)]$$

Component	x	Variables				
For ozone	x=ml	l is ozone content in atmosphere in cm				
For water vapor	x=mw	W is atmospheric perceptible water or columnar water vapor in gcm-2				
For aerosol	x=mb	Atmospheric visibility of haziness based on given aerosol type and size distribution				
For Other Gases	x=m	Relative optical air mass value				
Atmospheric variable	Ozone (/)	Water vapor ( <i>w</i> )	Aerosol (b)	Other gases		
a	0.0184	0.002	1.053	-5.4×10^(-5)		
b	0.0004	1.67 × 10^(-5)	0.083	-3.8×10^(-6)		
С	0.022	0.094	0.3345	0.0099		
d	-0.66	-0.693	-0.668	-0.62		
			Sou	rce: Paulescu and Schlett (2003)		



Diffuse Radiation

$$l_{diff} = S_R + S_A + S_M$$

S<sub>R</sub> Rayleigh Scattering
S<sub>A</sub> Aerosol scattering
S<sub>M</sub> Multiple scattering

#### Rayleigh Scattering

$$S_{R} = 0.79I_{0}\tau_{oz}\tau_{g}\tau_{w}(1 - (1 - w_{0})(1 - m + m^{1.06})(1 - \tau_{A}))\frac{0.5(1 - \tau_{R})}{1 - m + m^{1.06}}$$

#### Aerosol scattering

$$S_{A} = 0.79 s_{0} \tau_{oz} \tau_{g} \tau_{w} \left( \frac{\tau_{A}}{1 - (1 - w_{0})(1 - m + m^{1.06})(1 - \tau_{A})} \right) \times \frac{f_{c}}{1 - m + m^{10.6}} \left( 1 - \frac{\tau_{A}}{(1 - (1 - w_{0})(1 - m + m^{1.06})(1 - \tau_{A})\tau_{R})} \right)$$

#### >Multiple scattering

$$S_M = (S_{dir} + S_R + S_A) \frac{\rho_g \rho_a}{1 - \rho_g \rho_a}$$

 $\mathcal{P}_g$  is ground albedo and  $\mathcal{P}_a$  atmospheric albedo respectively

$$\rho_a = 0.0685 + (1 - f_c)(1 - \frac{\tau_A}{1 - (1 - w_0)(1 - m + m^{1.06})})$$

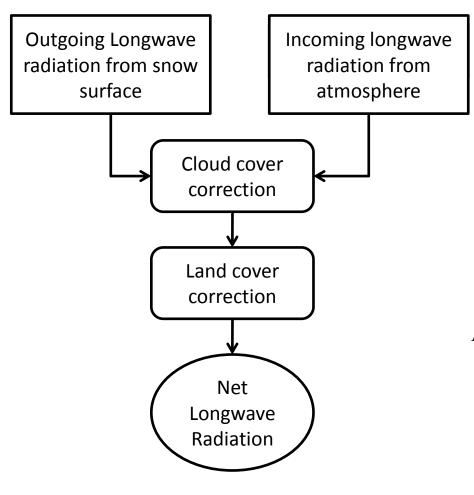
 $\mathcal{W}_0$  is single scattering albedo assumed as 1

 $f_c$  is ratio of forward to backward scattering calculated as  $f_c = 0.9302 \cos(\theta_s)^2$ 





#### **Longwave Radiation**



**Incoming Longwave Radiation** 

$$LW_{in} = \sigma * \varepsilon_{air} * T_{air}^{4}$$

 $\sigma$  Stefan Boltzmann Constant (5.67×10<sup>-8</sup> Wm<sup>-2</sup>K<sup>-4</sup>)

- $\mathcal{E}_{air}$  Emissivity of air
- $T_{air}$  Air Temperature

#### **Outgoing Longwave Radiation**

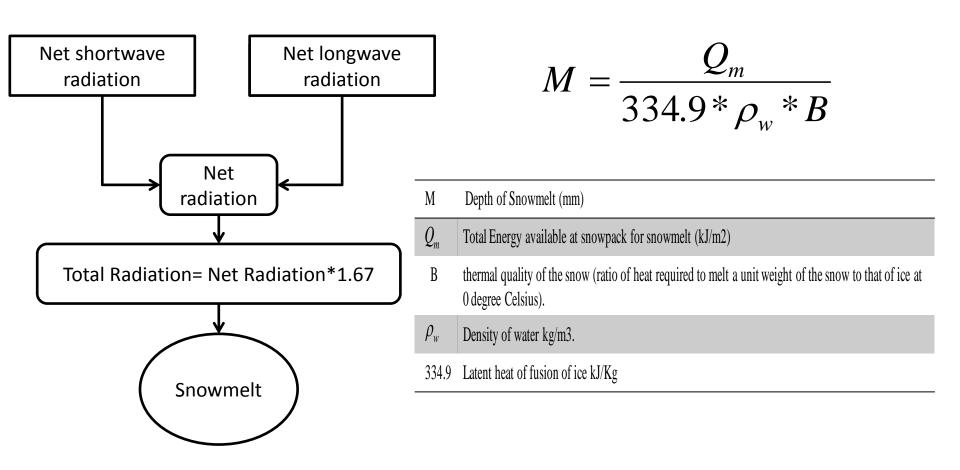
$$LW_{out} = \sigma * \varepsilon_{snow} * T_{lst}^{4} + LW_{in}(1 - \varepsilon_{snow})$$

 $\sigma$  Stefan Boltzmann Constant (5.67×10<sup>-8</sup> Wm<sup>-2</sup>K<sup>-4</sup>)

- $\mathcal{E}_{snow}$  Emissivity of snow
- *T*<sub>*lst*</sub> Snow Surface Temperature



### **Methodology for Snowmelt Calculation**





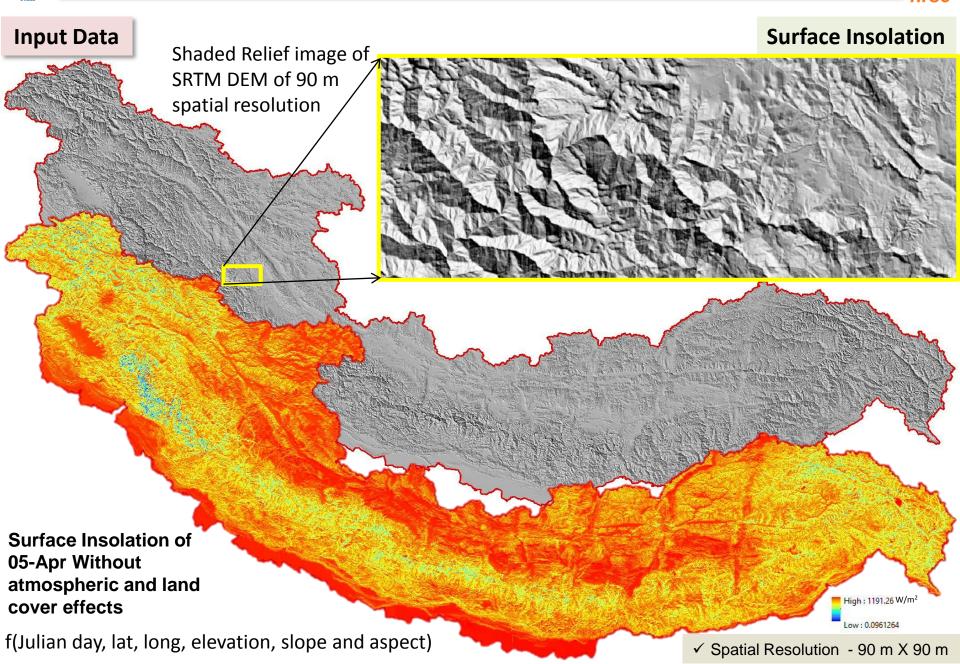


#### Input Data

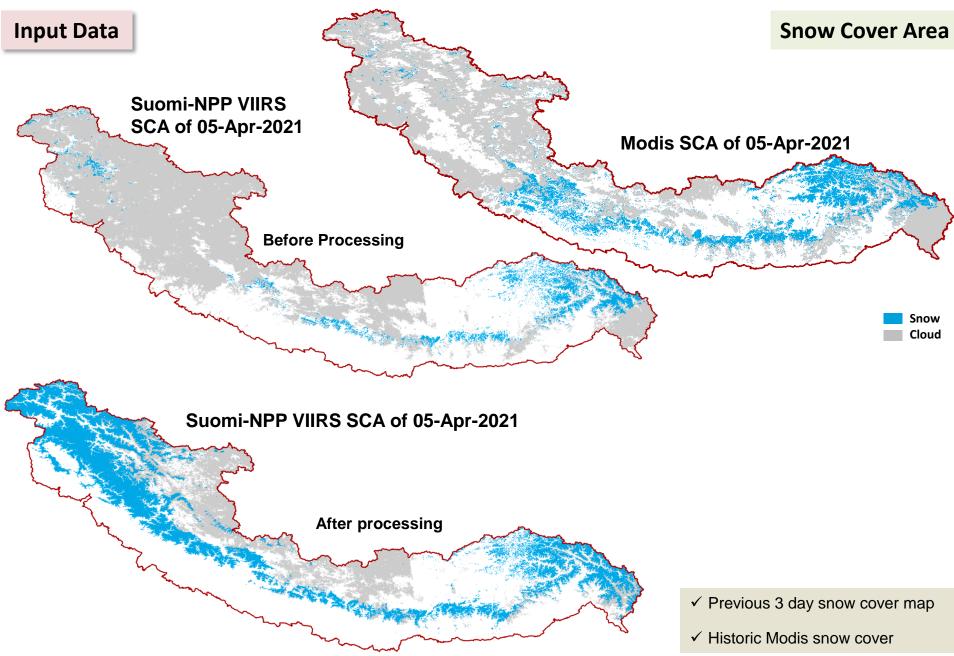
Input Data	Spatial Resolution	Source	
Insolation – f(Julian day, lat, long, elevation, slope and aspect)	90 m	SRTM DEM	
Snow Cover Area, Snow Albedo, LST	1 km	Suomi-NPP VIIRS satellite data - daily	
AOD	11 km	INSAT-3D/3DR Imager – Half hourly	
Water Vapour	11 km	INSAT-3D/3DR Sounder– Hourly	
Ozone	11 km	INSAT-3D/3DR Sounder– Hourly	
Cloud Mask	4 km	INSAT-3D/3DR Imager – Half hourly	
GEFS - Rainfall and Air Temperature forecast	48 km	Daily	
Land Use /Land Cover Map	56 m	Resourcesat-2 AWiFS satellite data	
Soil Map		NBSS&LUP and HWSD	
Glacier Map		Prepared from GLIMS, ICIMOD, and RGI Glacier Maps	



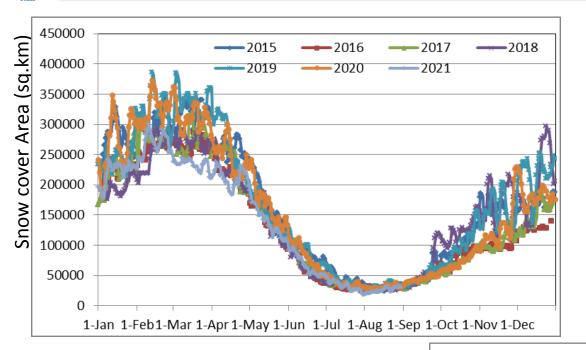








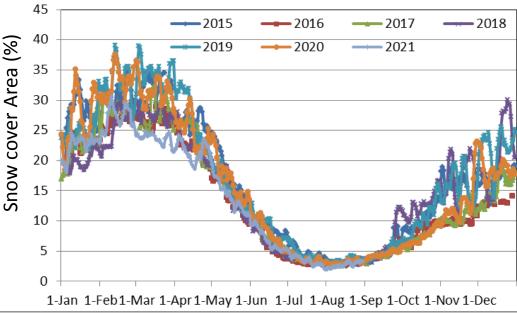




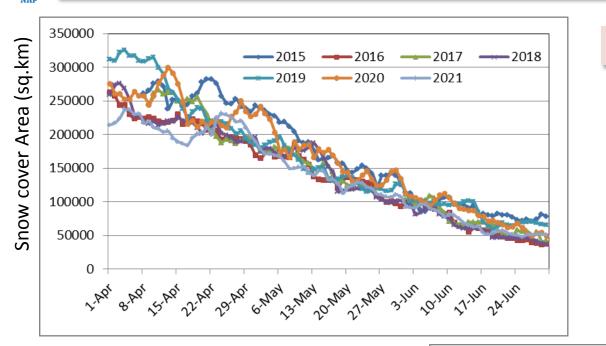
#### Variations of Snow Cover Extent

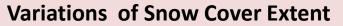
#### Snow Cover Area Information

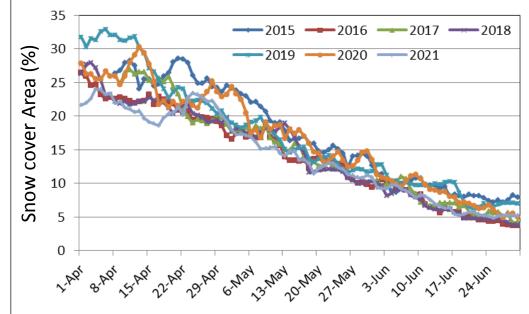
Year	Max	Min	Max	Min
2015	3,48,713	71,156	35.23	7.19
2016	2,86,035	26,728	28.90	2.70
2017	3,19,363	40,465	32.27	4.09
2018	2,96,539	24,451	29.96	2.47
2019	3,86,127	59,613	39.02	6.02
2020	3,72,313	26,425	37.62	2.67
2021	3,00,625	18,720	30.38	1.85



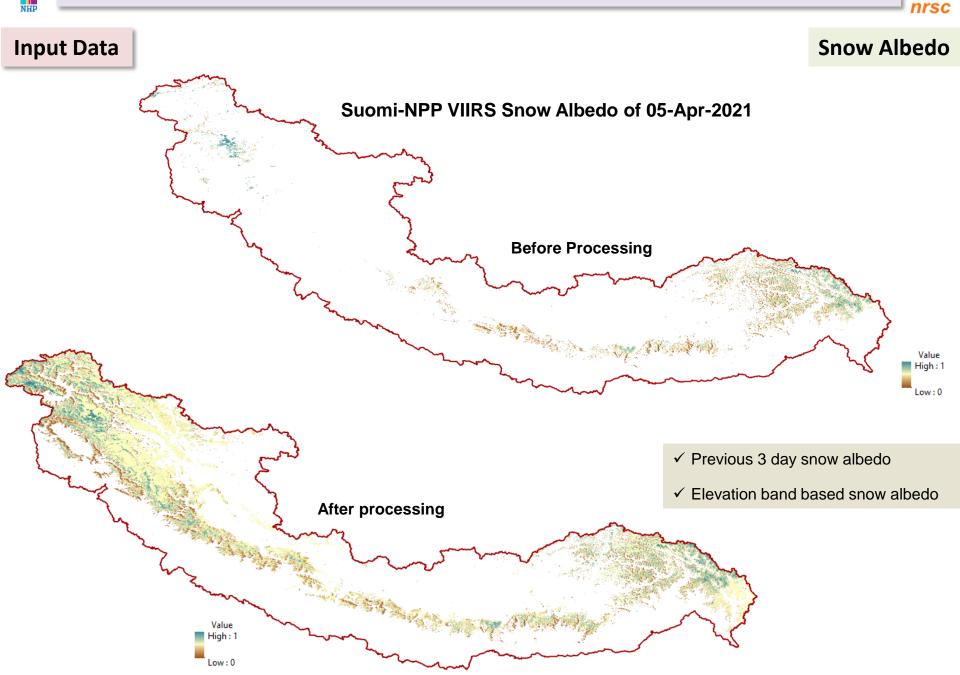








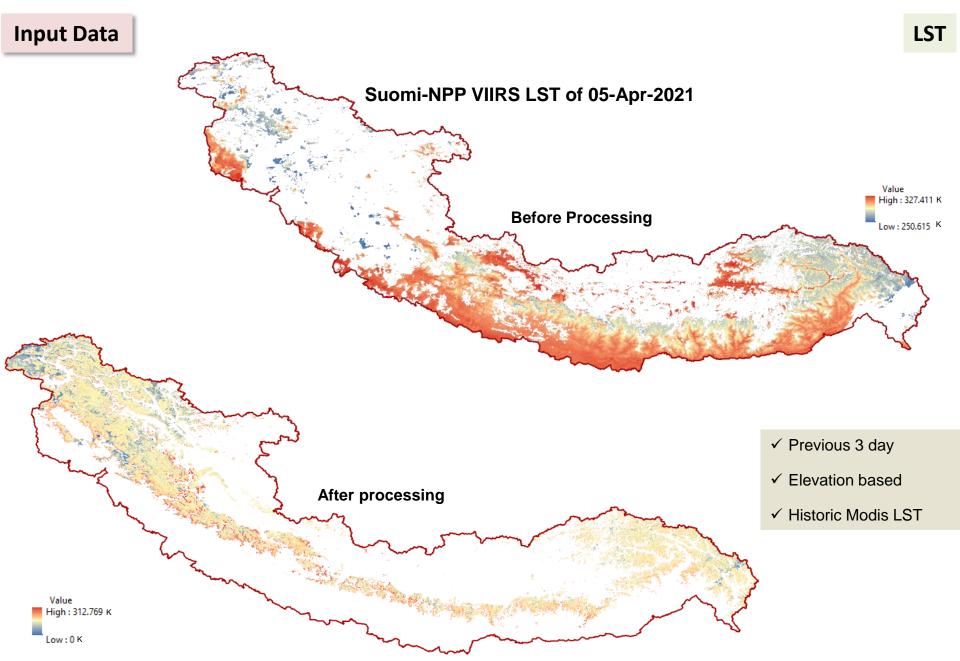
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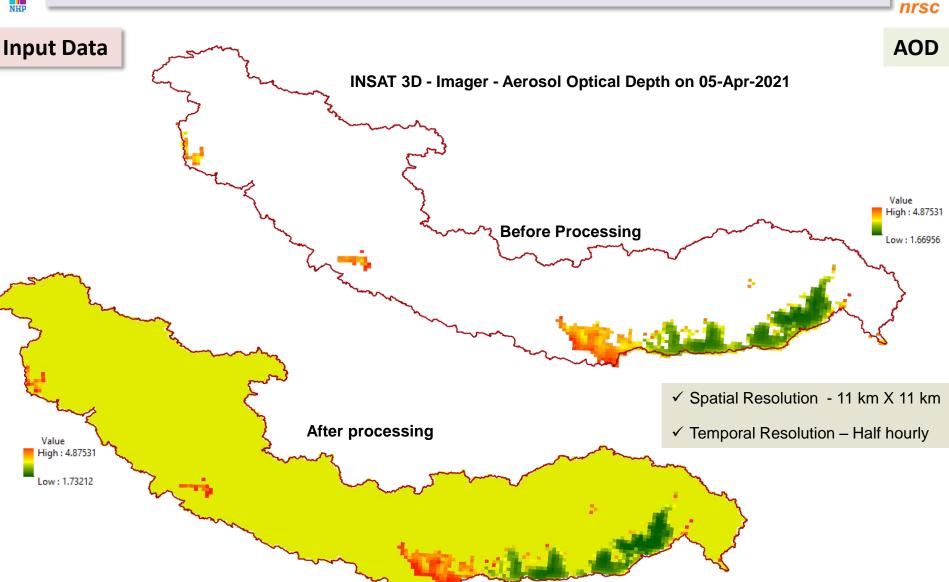
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## **Snow Cover Dynamics and Snowmelt Runoff Products & Services**

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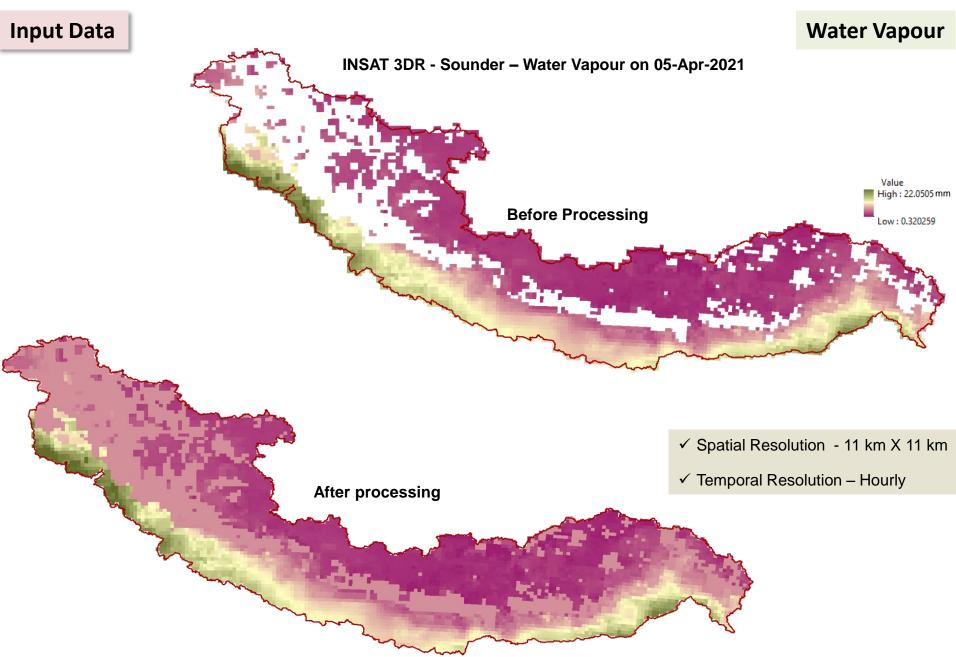
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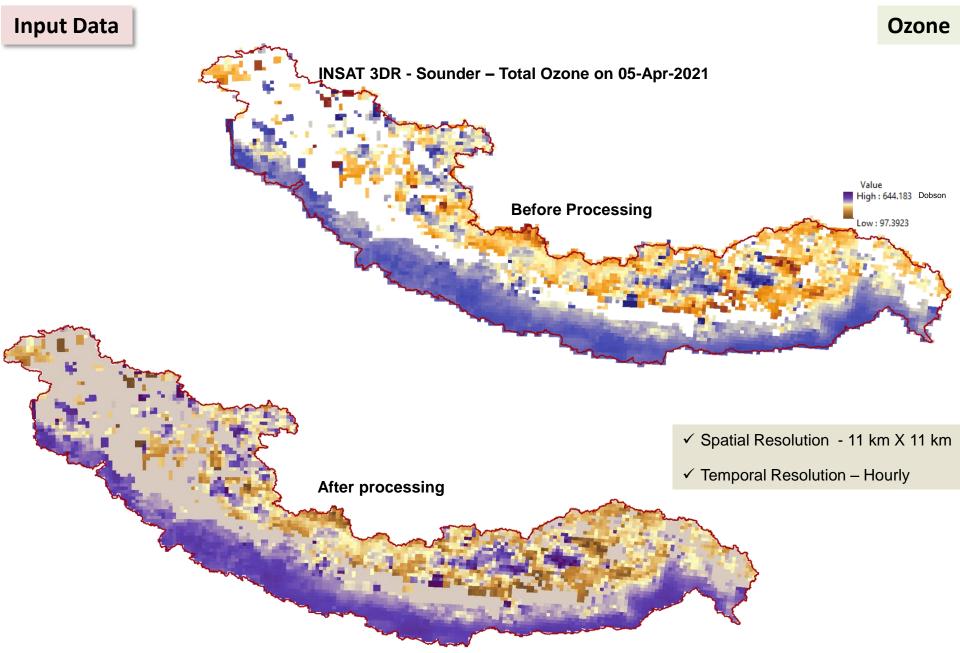
$$τ_{a(}(λ) = βλ^{-α}$$

 $\tau_{a(}(\lambda)$  = aerosol optical depth or thickness at wavelength  $\lambda$  (in  $\mu$ m)  $\beta$  = Angstrom turbidity coefficient is related to aerosols content  $\alpha$  = Angstrom wavelength exponent is related to the size distribution of particles

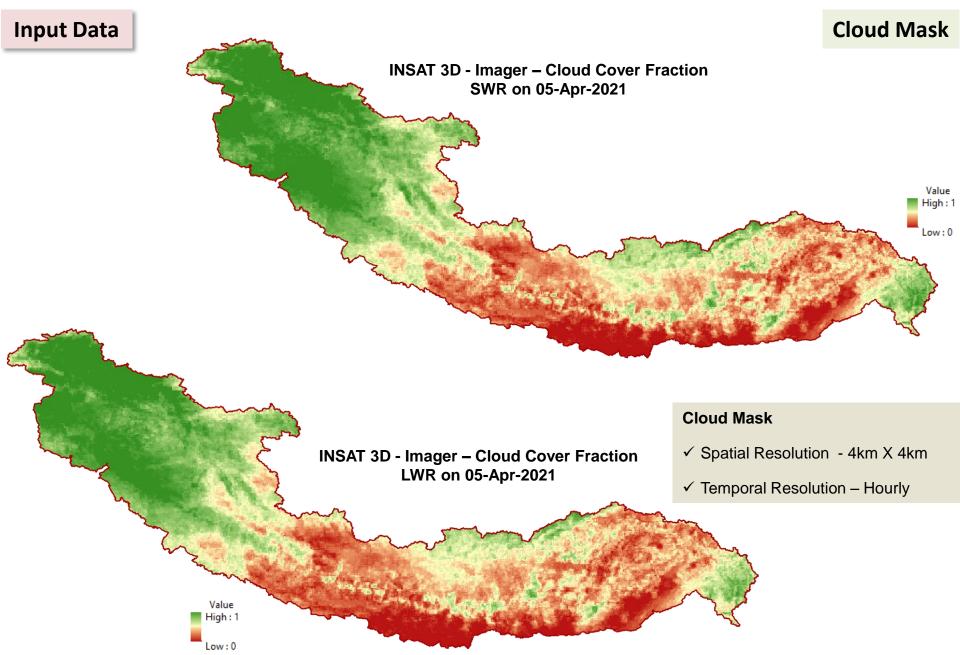




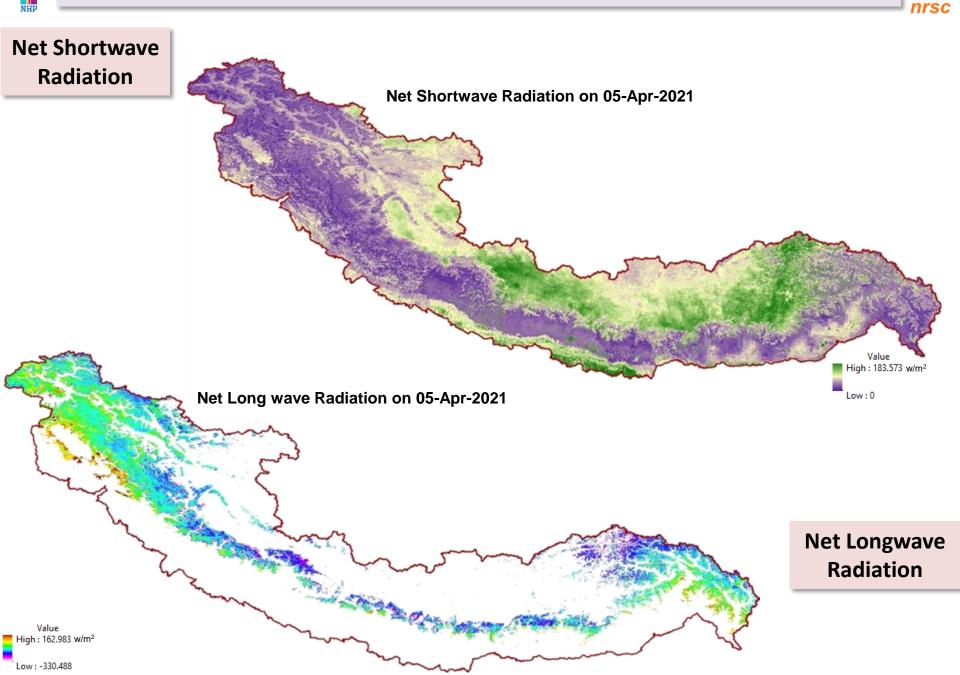








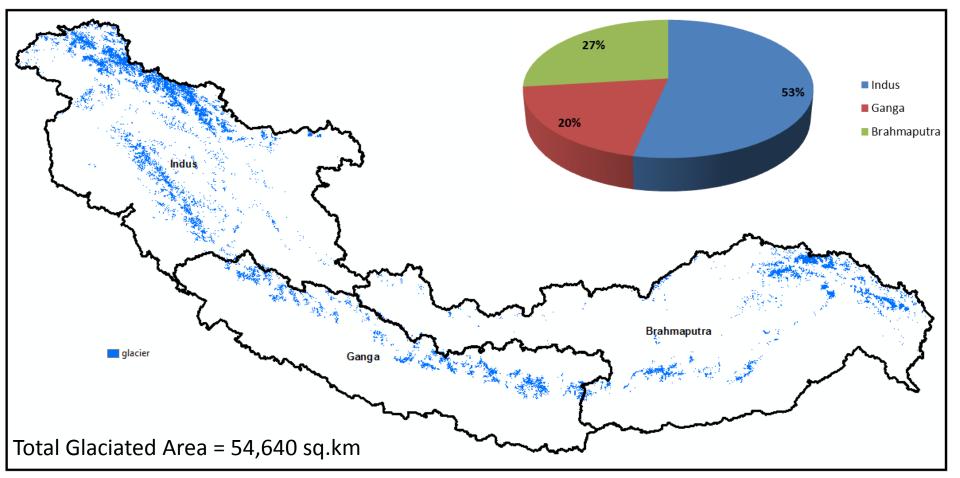
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#### Glaciers

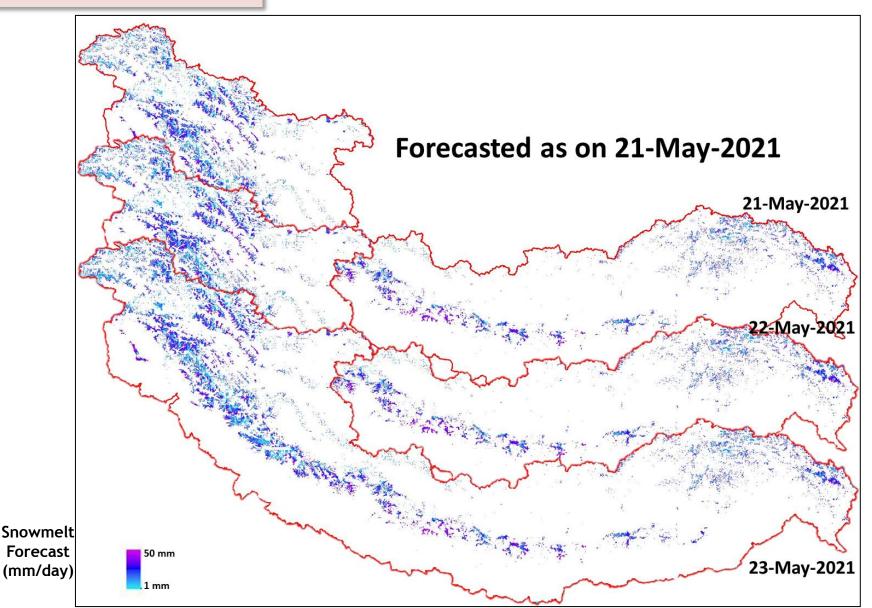
Glaciers mapped from GLIMS, ICIMOD, and RGI for Indian Himalayas region



Global Land Ice Measurements from Space (GLIMS) International Centre for Integrated Mountain Development (ICIMOD) Randolph Glacier Inventory (RGI) version 6  ✓ Glacier melt computed using empirical approach

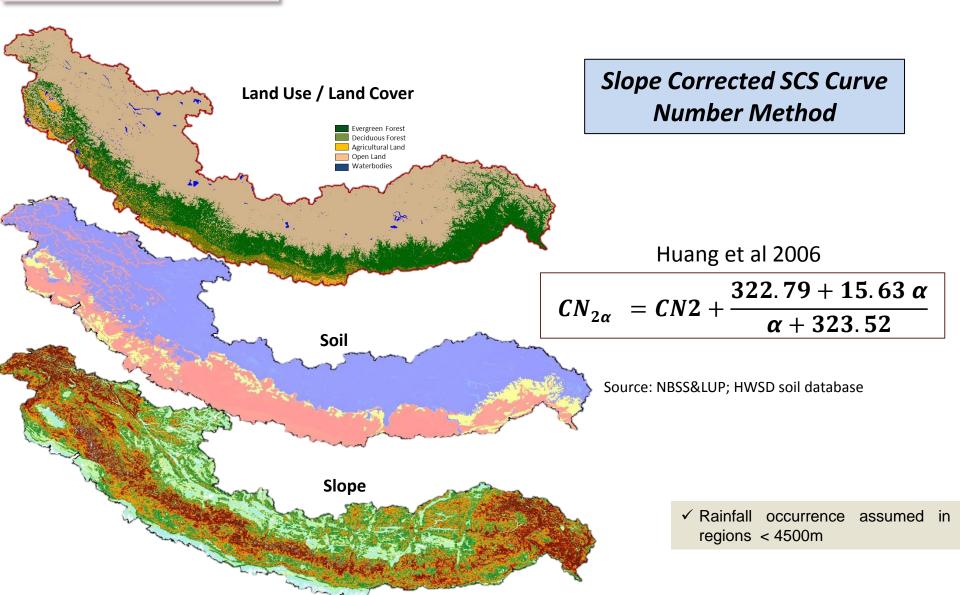


#### **Snowmelt Runoff Forecast**





#### **Rainfall-Runoff Modelling**









## Approach for 3-day Forecasting

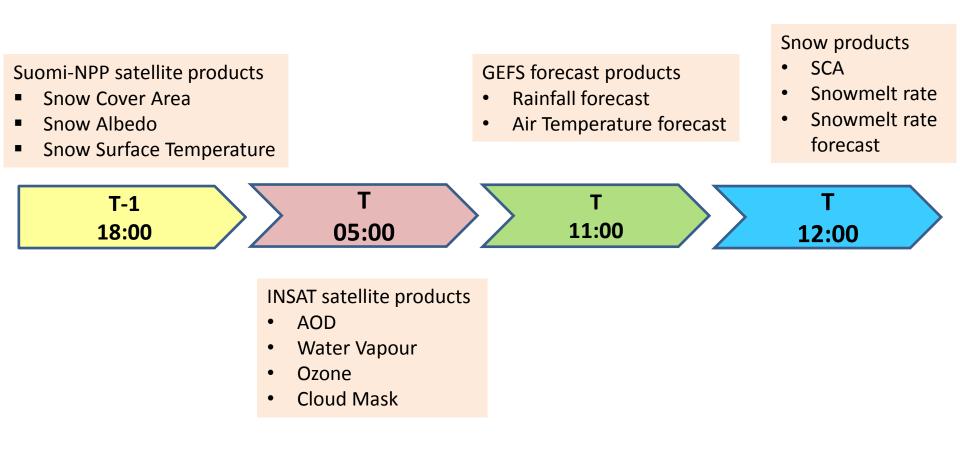
#### Adopted the following approach for inputs

- 1. All snow inputs (SCA, SA, LST, AOD, Ozone, WV) of previous day is considered as same for next 3days of forecast
- 2. Rainfall forecast for next 3 days
- 3. Air Temperature forecast for next 3 days











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#### **Assumptions and Limitations**

- Net SW Radiation and net LW Radiation estimated, out of all (7) EB components. For other EB components, it is assumed that they constitute about 40% of total energy input based on literature
- Empirical formula used for estimation of cloud cover effects on SW and LW radiation
- Suitable coefficients for Land cover classes are assumed for estimation of SW and LW radiation
- Empirical relationship was used to estimate air emissivity as a function of air temperature.
- It is assumed that rain fed area in Himalayan basins is generally below 4,500 m in elevation.
- Glacier melt is assumed to occur from middle of May. Glacier melt is estimated assuming suitable runoff coefficients considering the total glacier area within the basin. However, the level of exposure of glaciers may vary temporally and spatially. Year to year this exposure may differ.
- Even though cloud cover effect on net radiation, SCA, SA, LST is minimised, it may vary in actual conditions
- LST captured at 1030 Hrs is assumed to represent mean temperature of the day.
- Non-availability of snow depth and SWE data at fine resolution is major limitation
- Observed discharge at 0800 Hrs is assumed to represent that day's average discharge at that station. This assumption may not be valid as the discharge varies during the day.



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#### Utility of Snow Products

- Irrigation, Hydropower & Drinking water allocations
- Planning and Management of Reservoir Operations
- Planning of new hydropower projects
- Essential Climate Variable (ECV) for climate change studies
- Planning of new transport network

# THANK YOU